

Dhofar 378

Basalt
15 grams

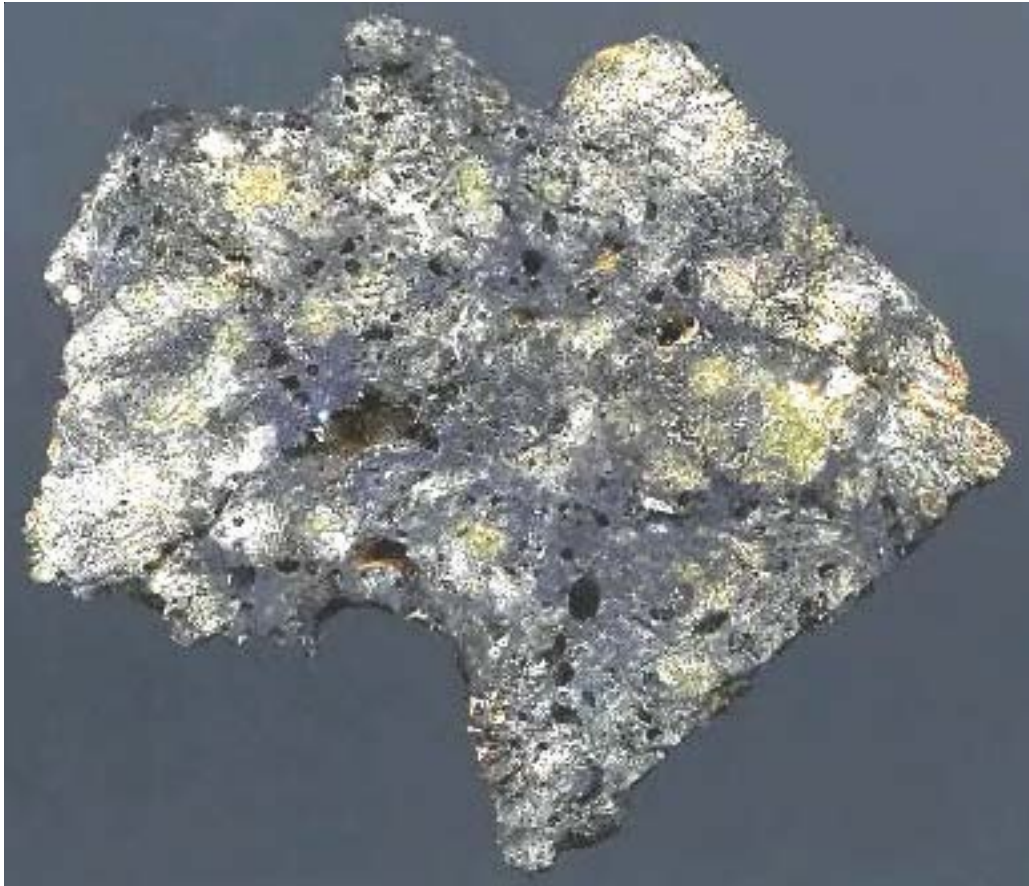


Figure 1: Photo of piece of Dhofar 378 (found on internet).

Introduction

Ikeda *et al.* (2002) first reported another Martian meteorite (Dho 378) from Oman (found in an area near Dho 019). Dho 378 is somewhat similar to Los Angeles, except that the plagioclase has been shocked to Plag-glass, instead of maskelynite (Ikeda *et al.* 2006). Dreibus *et al.* (2002) found that Dho 378 is enriched in Na, Sr and feldspar compared with other shergottites. An age of formation of 157 m.y. has been determined.

Petrography

Dho 378 is reported to have “a doleritic or microgabbroic texture, and grain sizes of the main minerals about 1 mm” (Ikeda *et al.* 2002). It has a fresh black fusion crust and has been highly shocked.

Dho378 is a basaltic shergottite made up of about equal amounts of clinopyroxene and plagioclase (now glass)

with minor mesostasis including alkali feldspar, silica, K-rich glass, Ca-phosphates and regions of hedenbergite, pyroxferroite, fayalite and pyrrhotite (Ikeda *et al.* 2006).

Vesicular shock-melt veins are common, sometimes cutting through pyroxene and Plag-glass grains.

Modal Mineralogy for Dhofar 378

	Ikeda <i>et al.</i> (2006)
Pyroxene	49.3%
Plagioclase (glass)	43.5
Opaques	2
Phosphates	1.4
Fayalite	tr.
Silica	tr.
Sulfide	0.6
Mesostasis	3.2

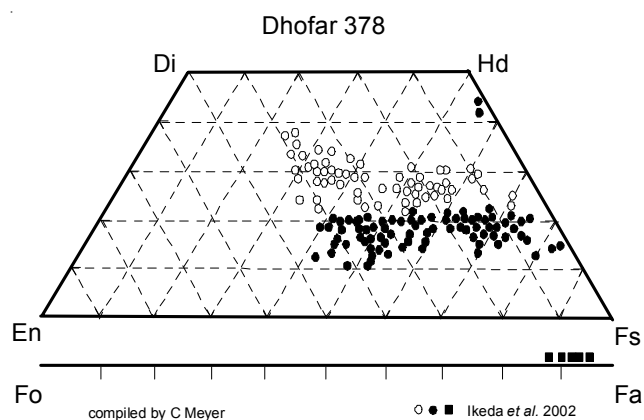


Figure 2: Pyroxene and olivine composition diagram for Dhofar 378 (replotted from Ikeda *et al.* 2002).

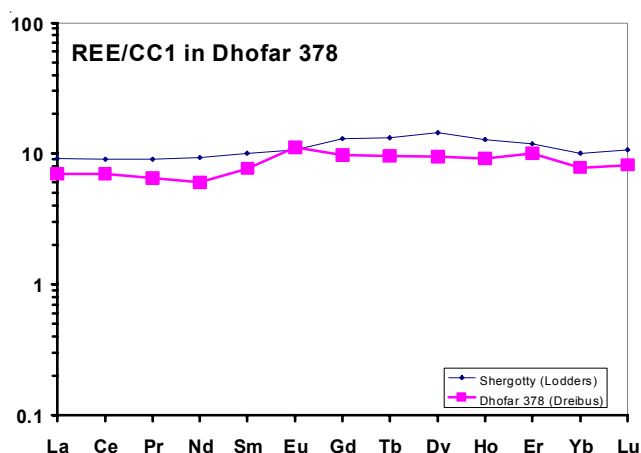


Figure 3: Rare earth element diagram for Dhofar 378 compared with Shergotty (data from Dreibus *et al.* 2002 and Lodders 1998).

Mineral Chemistry

Pyroxene: Pyroxene shows undulatory extinction and mosaicism due to shock (Mikouchi and McKay 2006). Subcalcic clinopyroxene is exsolved and chemically zoned (figure 2). The Fe/(Mg+Fe) ratio varies from 0.4 to 0.9 (Ikeda *et al.* 2002). Both Fe-rich hedenbergite and pyroxferroite are present.

Plagioclase: Plagioclase (An₃₃₋₅₃) has been shock-melted to glass, which has re-crystallized at the rims (Ikeda *et al.* 2006). Large plagioclase areas (probably several grains) up to 2 x 3 mm in size contain vesicles (~500 microns) with “dirty halos” around them.

Opaques: Titanomagnetite includes ilmenite lamellae (Ikeda *et al.* 2006).

Phosphates: Both whitlockite (merrillite) and minor apatite are present in the mesostasis of some thin sections. Compositions have been determined by Ikeda *et al.* (2006).

Calcite: Calcite occurs as a weathering product (Ikeda *et al.* 2006).

Silica: Both quartz and tridymite have been identified by Raman spectra (Ikeda *et al.* 2006).

Pyroxferroite: Small grains of pyroxferroite have been identified by Raman spectra (Ikeda *et al.* 2006). In Dho378 it has apparently formed as a secondary phase after the shock event.

Whole-rock Composition

Dreibus *et al.* (2002) and Ikeda *et al.* (2006) have analyzed Dho 378 (table 1). They explain a slight positive Eu anomaly as due to high plagioclase/whitlockite ratio in the small split they analyzed (figure 3). They note that Dho 378 is “weakly contaminated with terrestrial U and K”, as is the case for all Martian meteorites collected from “hot desert” sites.

Radiometric Age Dating

Nyquist *et al.* (2006) determined a crystallization age of 157 ± 24 m.y. by the Sm-Nd isochron method (figure 4). Park and Bogard (2006) studied the isotopic release of Ar from plagioclase and concluded that a major shock event occurred at ~ 143 m.y. ago (figure 5). The Rb-Sr isotopic system is highly disturbed by terrestrial contamination (calichi). However, the initial Sr isotopic ratio was found to be like that of Los Angeles (Nyquist *et al.* 2006).

Other Isotopes

The oxygen isotope composition, determined by Mayeda and Clayton and reported by Ikeda *et al.* (2002), is $\delta^{18}\text{O} = +4.46\text{‰}$ and $\delta^{17}\text{O} = +2.52\text{‰}$, indicating that this meteorite is Martian.

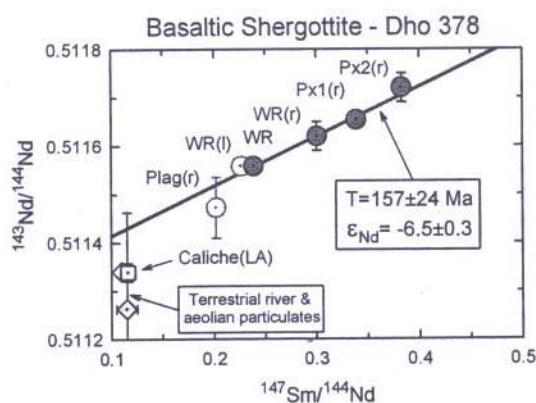


Figure 4: Sm-Nd isochron diagram for Dhofar 378 (from Nyquist et al. 2006).

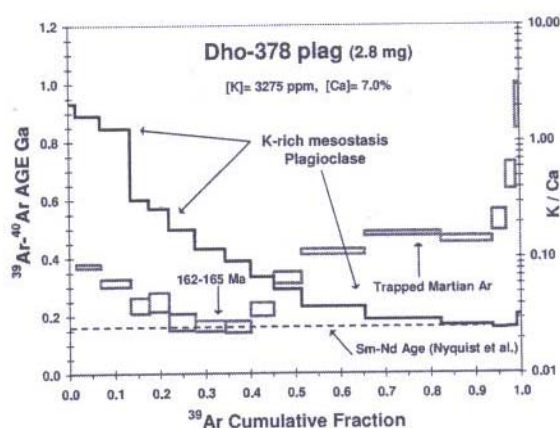


Figure 5: Ar plateau diagram for Dhofar 378 (but without the plateau)(from Park and Bogard 2006).

Table 1: Composition Dhofar 378

reference	Dreibus2002	Ikeda 2002	Ikeda2006	Ikeda2006
weight		fusion	crust	550 mg
SiO ₂		48.08	(b)	49.88 (c) 49 (d)
TiO ₂		1.18	(b)	0.98 (c) 1
Al ₂ O ₃		9.5	(b)	10.08 (c) 15.8
FeO	15.65	(a)	21.11	(b) 19.94 (c) 15.66
MnO	0.38	(a)	0.55	(b) 0.48 (c) 0.38
CaO	10.46	(a)	9.76	(b) 10.32 (c) 10.45
MgO			5.35	(b) 5.66 (c)
Na ₂ O	2.6	(a)	2.31	(b) 1.98 (c) 2.64
K ₂ O	0.2	(a)	0.15	(b) 0.17 (c) 0.2
P ₂ O ₅			0.9	(b) 0.7 (c) 1.2
sum		98.89		100.23 (c)
Li ppm				
Be				
F				
S				
Cl				
Sc	43.7	(a)		
V				
Cr	260	(a)		274 (c)
Co	29.3	(a)		
Ni	<40	(a)		
Cu				
Zn	77	(a)		
Ga	23.6	(a)		
Ge				
As				
Se				
Br	0.89	(a)		
Rb				
Sr	120	(a)		
Y				
Zr				
Nb				
Mo				
Pd ppb				
Ag ppb				
Cd ppb				
In ppb				
Sb ppb				
Te ppb				
Cs ppm	0.4	(a)		
Ba	33	(a)		
La	1.64	(a)		
Ce	4.17	(a)		
Pr				
Nd	2.7	(a)		
Sm	1.13	(a)		
Eu	0.627	(a)		
Gd	1.9	(a)		
Tb	0.35	(a)		
Dy	2.3	(a)		
Ho	0.51	(a)		
Er				
Tm				
Yb	1.27	(a)		
Lu	0.2	(a)		
Hf	1.49	(a)		
Ta	0.14	(a)		
W ppb				
Re ppb				
Os ppb				
Ir ppb	<6	(a)		
Au ppb	7.7	(a)		
Tl ppb				
Bi ppb				
Th ppm	0.3	(a)		
U ppm	0.1	(a)		

technique: (a) INAA; (b) elec. Probe, (c) XRF, (d) INAA

References for Dhofar 378

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